

Sediment Production in Tropical Catchments: An Integration of USLE and Index of Connectivity in Antonina, Southern Brazil

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Keywords

Hydrosedimentological Connectivity
Soil Loss
Sedimentation

Abstract

This study analyzes sediment production and export in basins draining into Antonina Bay (PR), using the Universal Soil Loss Equation (USLE), the Connectivity Index (CI), and the Sediment Delivery Rate (SDR) to calculate an effective sediment export index (Eeffective). The study region encompasses the Cachoeira, Pequeno, Cacatu, and Faisqueira river basins, located in a fluvio-marine dam area, where agricultural use is concentrated on alluvial deposits. The results indicate that the USLE predicts average losses ranging from 1.29 to 3.11 t ha⁻¹ yr⁻¹ per basin. In contrast, the Eeffective reduces losses by 70 to 72% (0.36 to 0.88 t ha⁻¹ yr⁻¹), considering retention and disconnectivity along the slopes. Low-elevation regions associated with agricultural use, particularly in the Lower Cachoeira River, are the main points of sediment production and export. The USLE, IC, and Eefetivo maps indicate that most of the transported sediment does not reach the outlet, highlighting the role of connectivity. Thus, combining estimates of potential erosion with connectivity metrics can improve the spatial representation of sediment export and guide conservation measures (no-till farming, agroforestry systems, and riparian forest restoration) to reduce coastal siltation and determine priority areas for restoration.

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INTRODUCTION

Sheet erosion is one of the most significant soil-degradation processes, characterized by the continuous and relatively uniform removal of surface particles by runoff (Bertoni *et al.*, 2012; Kopittke *et al.*, 2025). Although often underestimated due to its low visibility compared to gully or ravine erosion, its cumulative effects can lead to substantial losses of fertile soil, compromising agricultural productivity and environmental sustainability (Morgan, 2005; Poesen, 2018). Globally, sheet erosion is estimated to account for the loss of billions of tons of soil each year, with direct implications for food security and ecosystem functioning (Pimentel *et al.*, 1995; Borrelli *et al.*, 2017; Quinton; Fiener, 2023; Xiong; Leng, 2024).

Sheet erosion is particularly critical in tropical and subtropical regions, where intense rainfall promotes the removal of the most fertile surface horizons and ultimately reduces agricultural productivity. Additional impacts include siltation, environmental degradation, and increased restoration costs (Labrière *et al.*, 2015).

Soil-loss rates vary widely among catchments: approximately 15 t ha⁻¹ year⁻¹ in the Jacaré-Guaçu River catchment in São Paulo, southeastern Brazil (Souza, 2016), and more than 48 t ha⁻¹ year⁻¹ in the Juramento River catchment in Minas Gerais, also in southeastern Brazil (Oliveira; Leite, 2018). In agricultural areas with exposed temporary crops, losses can be even greater, reaching around 51.6 t ha⁻¹ year⁻¹, whereas under forest cover they decrease to only 3–4 kg ha⁻¹ year⁻¹ (Marinhascki, 2016). In Paraná (southern Brazil), estimates range from 10 t ha⁻¹ year⁻¹ (Souza *et al.*, 2018) to 15–20 t ha⁻¹ year⁻¹ in intensively managed areas (Krug, 2020), with hotspots exceeding 50–100 t ha⁻¹ year⁻¹ in exposed soils and access roads, and peak values up to 355 t ha⁻¹ year⁻¹ (Souza *et al.*, 2018). This spatial heterogeneity highlights the need for soil-conservation practices, vegetation restoration, and road-management strategies in critical areas (Krug, 2020; Demarchi, *et al.*, 2019; Xiong; Leng, 2024).

Beyond reducing soil fertility, sheet erosion contributes to the silting of rivers, lakes, and reservoirs, decreasing water availability and increasing dredging costs (ANA, 2020). Dredging is costly and causes significant environmental impacts in estuarine environments (Paula *et al.*, 2006), including hydraulic and sedimentological alterations such as changes in circulation patterns, water-column mixing, and increased turbidity (Liu *et*

al., 2010). Contaminant remobilization, such as trace metals and organic compounds, frequently occurs, affecting water quality and aquatic organisms (Monte *et al.*, 2019; Moreira *et al.*, 2021). Sediment removal also alters benthic habitats, reducing feeding and breeding areas and intensifying socio-environmental conflicts associated with dredging operations (Castro; Almeida, 2012).

The combined effects of siltation and repeated dredging emphasize the magnitude of the challenges faced in estuarine systems. In this context, the Paranaguá Estuarine Complex (PEC), located on the coast of Paraná, southern Brazil, stands out for its hydrodynamic, ecological, and geochemical complexity. Between 2009 and 2015, approximately 23.5×10^6 m³ of sediments were dredged in the PEC, requiring investments of 365.8 million reais by the Paranaguá and Antonina Port Authority (APPA – Associação de Portos de Paranaguá e Antonina) (Neto *et al.*, 2017).

Siltation in the area results from the accumulation of sediments derived from the catchments surrounding the Port Complex, produced by natural processes often intensified by inadequate land use (Rutyna *et al.*, 2021). This dynamic is particularly concerning given the increasing operational pressure on the port complex, which depends on maintaining navigability. In 2022, the PEC received 2,540 vessels, rising to 2,933 in 2023; that year, it handled 65.39 million tons of cargo and generated a net operating revenue of R\$ 620.96 million (Portos do Paraná, 2023; 2024). These conditions reinforce the importance of understanding sediment sources and the processes controlling sediment production and delivery to the system.

Understanding these processes is essential for guiding management actions, especially in systems under high logistical pressure such as the PEC. Modeling tools are fundamental for assessing basin susceptibility to sediment generation. Advances in computational modeling, such as the Universal Soil Loss Equation (USLE) and its variations (RUSLE, MUSLE), have enabled the prediction of high-risk areas and the optimization of management and restoration strategies (Renard *et al.*, 1997; Borrelli *et al.*, 2021; Kumar *et al.*, 2022). These models estimate potential sediment production but do not account for factors controlling sediment transfer and retention along watersheds.

Sediment moves through watersheds along slopes and channel networks (Bracken *et al.*, 2015). To address the limitations of traditional models and better represent sediment flow,

metrics such as the Index of Connectivity (IC) (Borselli *et al.*, 2008) have been used to evaluate the efficiency of transporting eroded material to the drainage network. The IC is directly related to the Sediment Delivery Ratio (SDR), which quantifies the fraction of sediments effectively delivered to watercourses (Ferro; Minacapilli, 1995), thereby allowing the estimation of effective sediment export (Eeffective), the portion of total erosion that reaches basin outlets (Vigiak *et al.*, 2012).

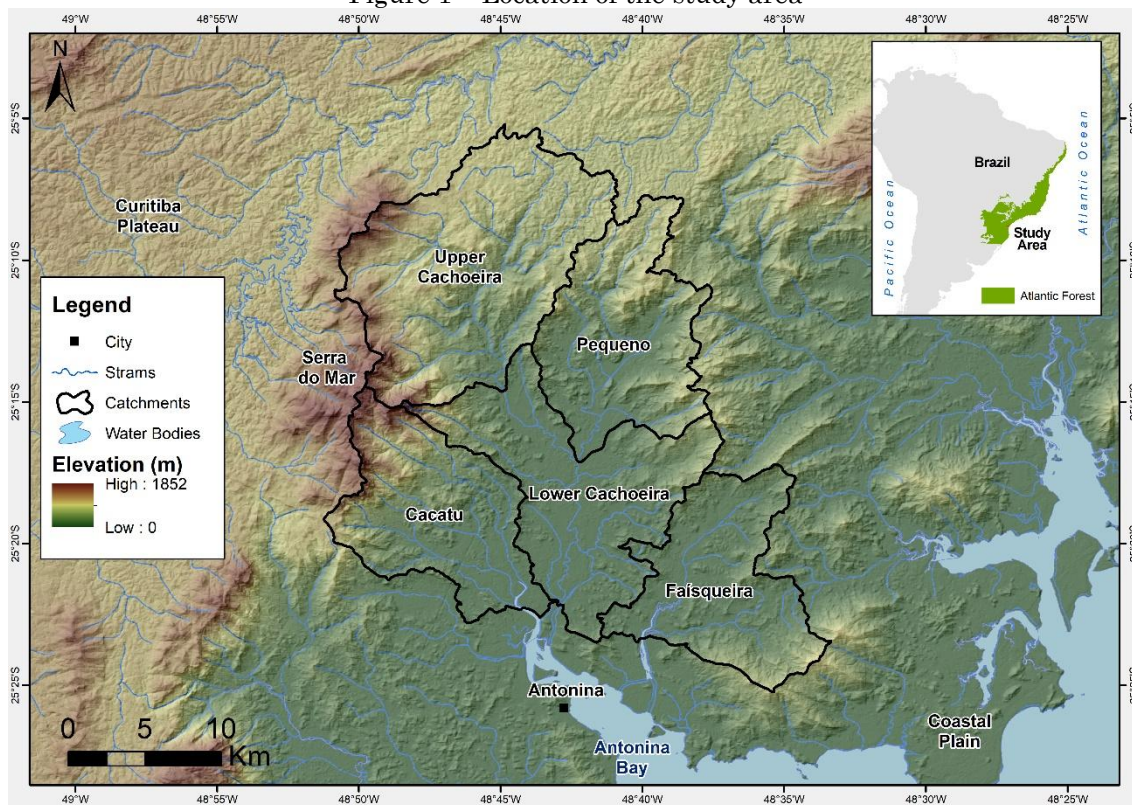
Given that total sediment production differs from the exported volumes, this article applies USLE, IC, SDR, and Eeffective to identify areas of highest sediment generation and export in the basins draining into Antonina Bay, southern Brazil, and to discuss the factors that intensify these processes in the region.

METHODS

Study Area

The study area comprises the catchments of the Cachoeira (upper and lower), Pequeno, Cacatu, and Faísqueira rivers, located in the municipality of Antonina, State of Paraná, southern Brazil (Figure 1). These basins are adjacent to the PEC and form part of the Degraded Areas Recovery Program (PRAD) within the Guaraqueçaba Environmental Protection Area. The program was implemented in the municipality of Antonina by the state-owned public company Paraná Ports as a requirement associated with dredging works carried out to deepen the navigation channel, access routes, and berths of the Port of Paranaguá.

Figure 1 – Location of the study area



Source: The authors (2025).

The studied drainage basins exhibit a strong predominance of forest cover, ranging from 88% to 95%, whereas agricultural land occupies relatively small areas, between 4% and 11%. Basin sizes vary from 103 to 180 km². Regarding average slope, the values show notable

differences among the basins: upper Cachoeira and Cacatu present the highest mean slopes (34% and 31%, respectively), while lower Cachoeira and Faísqueira display the lowest averages (18% and 20%) (Table 1).

Table 1 – Characteristics of the studied river basins

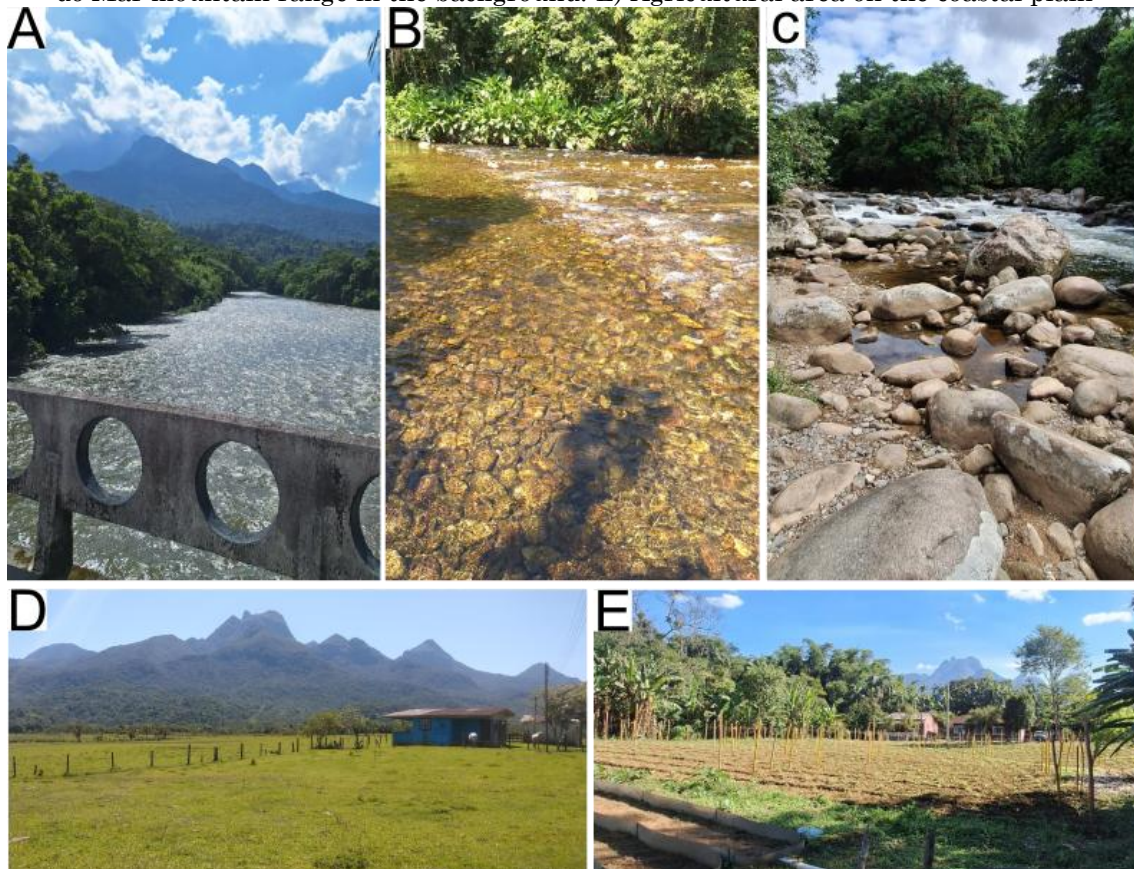
Catchment	Área (Km ²)	Land Use (%) *		Slope (%)
	Forest	Agriculture		Média (DP)
Upper Cachoeira	180	95	4	34,4 (23)
Lower Cachoeira	137	88	11	18,3 (18)
Cacatu	106	95	4	31,9 (28,3)
Faisqueira	103	95	4	20,9 (17,6)
Pequeno	112	93	7	30,5 (22,2)

Source: Souza *et al.*, 2020.

The region is characterized by vegetation formations typical of the Atlantic Forest, ranging from montane and high-montane forests in the Serra do Mar to pioneering fluviomarine and fluviolacustrine communities associated with the coastal plain (Roderjan, 2002; LAGEAMB, 2023). In the municipality of Antonina, approximately 87% of the territory is covered by forest vegetation (Souza *et al.*, 2020).

Among the anthropogenic land uses, diversified agriculture stands out, including rice, cassava, corn, banana, orange, and other crops (Figure 2). These activities are heavily concentrated in flat areas near rivers, where soils developed from fluvial deposits predominate (Ipardes, 2025).

Figure 2 –Landscapes of the study area. A) Bridge over the Cachoeira River, lower course. B) Pequeno River; C) Cachoeira River, upper course; D) Landscape of the coastal plain, with the Serra do Mar mountain range in the background. E) Agricultural area on the coastal plain



Source: The authors (2025).

The drainage area of Antonina Bay is predominantly underlain by high-grade metamorphic rocks, especially migmatites and gneisses, along with Archean and Proterozoic crystalline complexes. The region also includes the Guaratubinha Formation with its basic dikes, Tertiary deposits of the Alexandra

Formation, and extensive Holocene sediments, particularly alluvial deposits (Salamuni; Rocha, 2002; Ângulo, 2004).

The climate is classified as humid subtropical (Cfa) in the lowlands and humid mesothermal subtropical (Cfb) in the mountainous areas, with an average annual rainfall of 2,273 mm.

Summer is the wettest season, with monthly totals exceeding 300 mm, whereas winter precipitation generally remains below 150 mm (ClimateData, 2024; ITCG, 2008; Goudard; Paula, 2016).

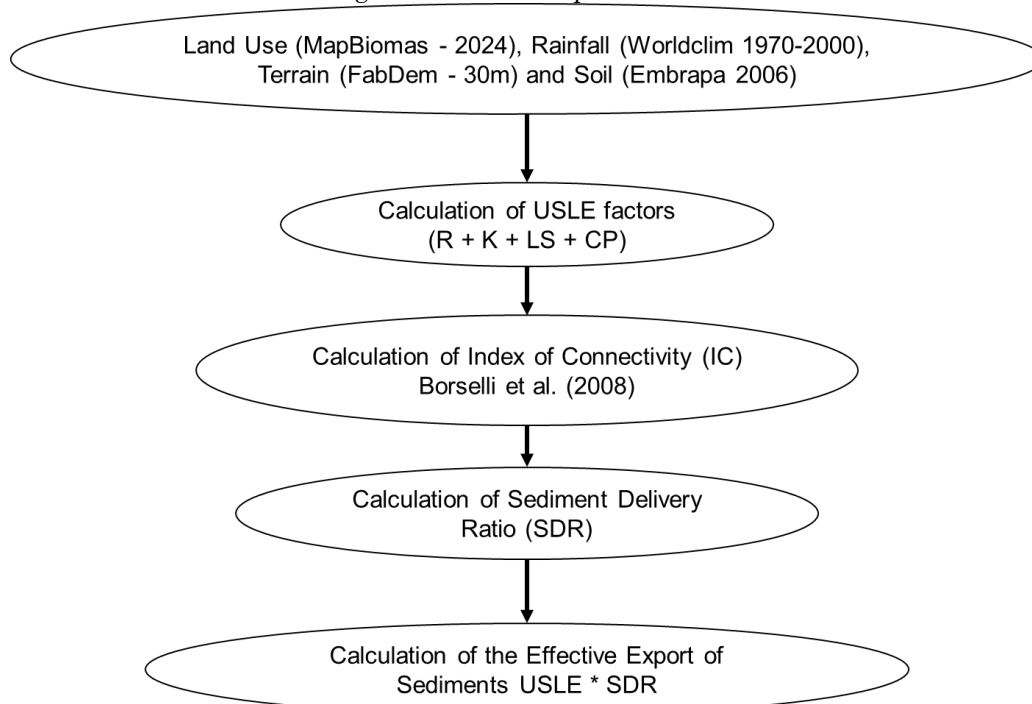
The predominant soils in the region are Acrisols, Cambisols, Gleysols, Leptosols and Fluvisols. Acrisols occur mainly in the less rugged portions of the Serra do Mar mountain range and in hilly areas, generally associated with Cambisols. These, in turn, are widely distributed in the Serra do Mar, and may occur in isolation or in association with other classes, especially in colluvial areas, hills, and mountains. Fluvisols are concentrated in valley bottoms, where they may occur alongside Gleysols, which also predominate in these environments, generally in isolation. Leptosols are characteristic of the mountainous portions of

the Serra do Mar, while Fluvisols develop in fluvial environments (Paula, 2010).

Estimation of Sediment Production

The analysis of sediment production and connectivity in the study area was carried out using multiple complementary approaches. Direct soil loss was estimated using the Universal Soil Loss Equation (USLE). Sediment connectivity was evaluated through the Index of Connectivity (IC). These two components were then integrated by applying the Sediment Delivery Rate (SDR), which enabled the estimation of the effective sediment export index (Eeffective) (Borselli *et al.*, 2008; Vigiak *et al.*, 2012; Cavalli *et al.*, 2013) (Figure 3).

Figure 3 – Work steps flowchart



Source: The authors (2025).

The USLE (Wischmeier; Smith, 1965; 1978) was applied to estimate potential sheet erosion in the study area. The USLE estimates average annual soil loss ($t\ ha^{-1}\ year^{-1}$) based on climatic, topographic, pedological, and land-management factors (Equation 1):

$$A = R \times K \times L \times S \times C \times P$$

Where A = average annual soil loss ($t\ ha^{-1}\ year^{-1}$)

R = rainfall erosivity factor ($MJ\ mm\ ha^{-1}\ h^{-1}\ year^{-1}$)

K = soil erodibility factor ($t\ ha\ h\ ha^{-1}\ MJ^{-1}\ mm^{-1}$)

LS = topographic factor (slope length and slope steepness)

C = land cover and management factor

P = conservation practices factor

Rainfall Erosivity – R Factor

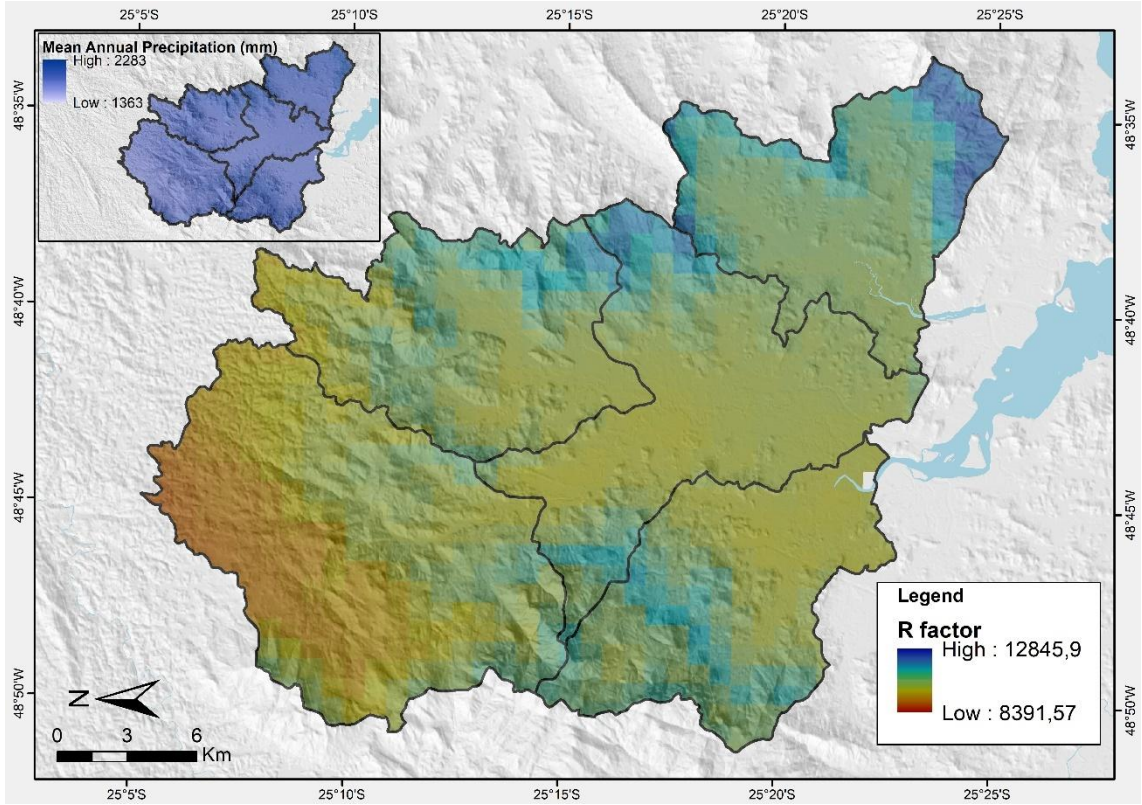
The R factor represents the erosive potential of rainfall (Figure 4), integrating both the intensity and duration of precipitation events. In this study, it was calculated using the equation proposed by Lombardi Neto and

Moldenhauer (1992), which estimates erosivity based on monthly and annual precipitation data (Equation 2):

$$EI = 68,73 \left[\left(\frac{p^2}{P} \right)^{0,841} \right]$$

Where *p* is the average monthly precipitation (mm) and
P is the average annual precipitation (mm).
The precipitation data were obtained from the WorldClim climatic database (Fick; Hijmans, 2017), with a spatial resolution of 1 km, corresponding to the 1970–2000 climatological normal.

Figure 4 – R-factor



Source: The authors (2025).

Soil Erodibility - K Factor

The K factor reflects the soil's susceptibility to erosion (Figure 5). This factor can be determined in the laboratory or by empirical equations (e.g., Wischmeier nomogram). However, for application in larger areas, it is common to assign K values to mapped soil units, such as

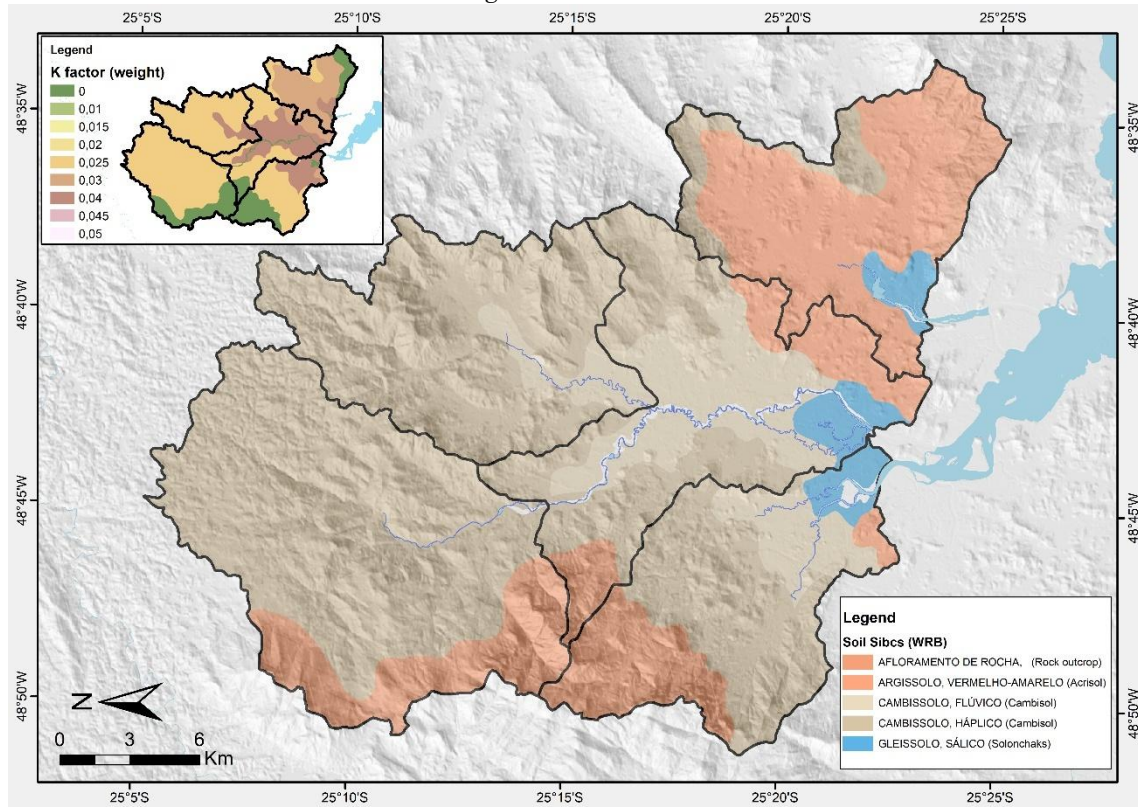
those of the Brazilian Soil Classification System (SiBCS). For this study, we used values adapted from the literature (Table 2) (Mello *et al.*, 2007; Coelho *et al.*, 2024). The soil data used were obtained from the work of Bhering *et al.*, (2007) at a scale of 1:250.000.

Table 2 – Soil classes found in the area and K values

Brazilian soil classification (Santos <i>et al.</i> , 2025)	Soil Class (WRB, 2022)	Estimated K (t·ha·h)/(MJ·mm·ha)	Weight
Rock outcrop	Rock outcrop	0,000	
Argissolo Vermelho-Amarelo	Acrisol	0,020 – 0,040	0,030
Cambissolo Háplico	Cambisol	0,020 – 0,035	0,025
Cambissolo Flúvico	Cambisol	0,030 – 0,050	0,040
Gleissolo Sáfico	Solonchaks	0,030 – 0,050	0,040

Source: Adapted from Coelho *et al.*, (2024).

Figure 5 –K Factor



Source: Bhering *et al.*, (2007).

Length/slope – LS Factor

The LS factor (Figure 6) represents the combined influence of slope length and gradient on soil erosion. This factor quantifies how topographic characteristics control erosion potential and is calculated through a mathematical formulation that integrates terrain geometry (Desmet; Govers, 1996).

The Digital Elevation Model (DEM) used to derive the LS factor was obtained from the FABDEM project (Forest and Buildings Removed from the Copernicus DEM). FABDEM is derived from the global Copernicus DEM and provides elevation data at 30 m spatial

resolution, delivered on a 1-arc-second grid (approximately 30 m at the equator) (Hawker *et al.*, 2022).

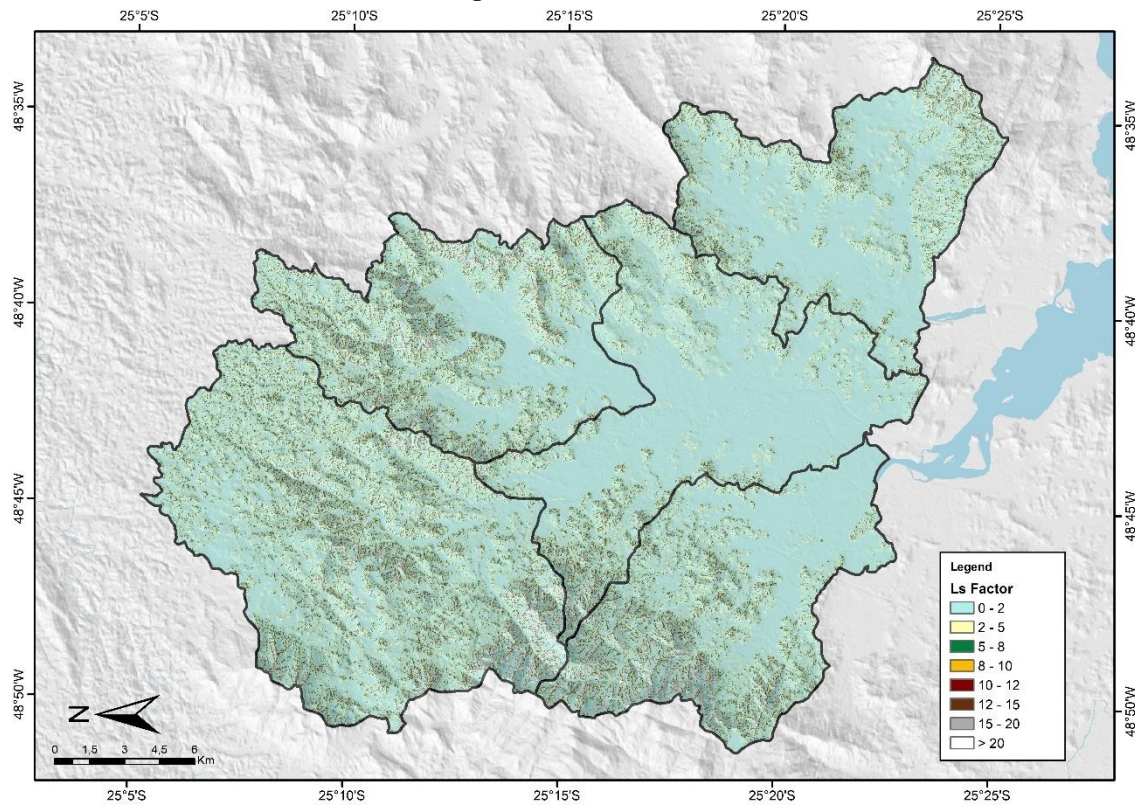
Reference values of 22.1 m for slope length and 0.09 (equivalent to approximately a 9% slope) were adopted in the calculation (Equation 3).

$$LS = \left(\frac{FA * 30}{22.1} \right)^{0.6} * \left(\frac{\sin(S * 0.01745)}{0.09} \right)^{1.3}$$

Where AF = Flow accumulation

S = Slope

Figure 6 – LS Factor



Source: The authors (2025).

Land Use and conservation practices – CP Factor

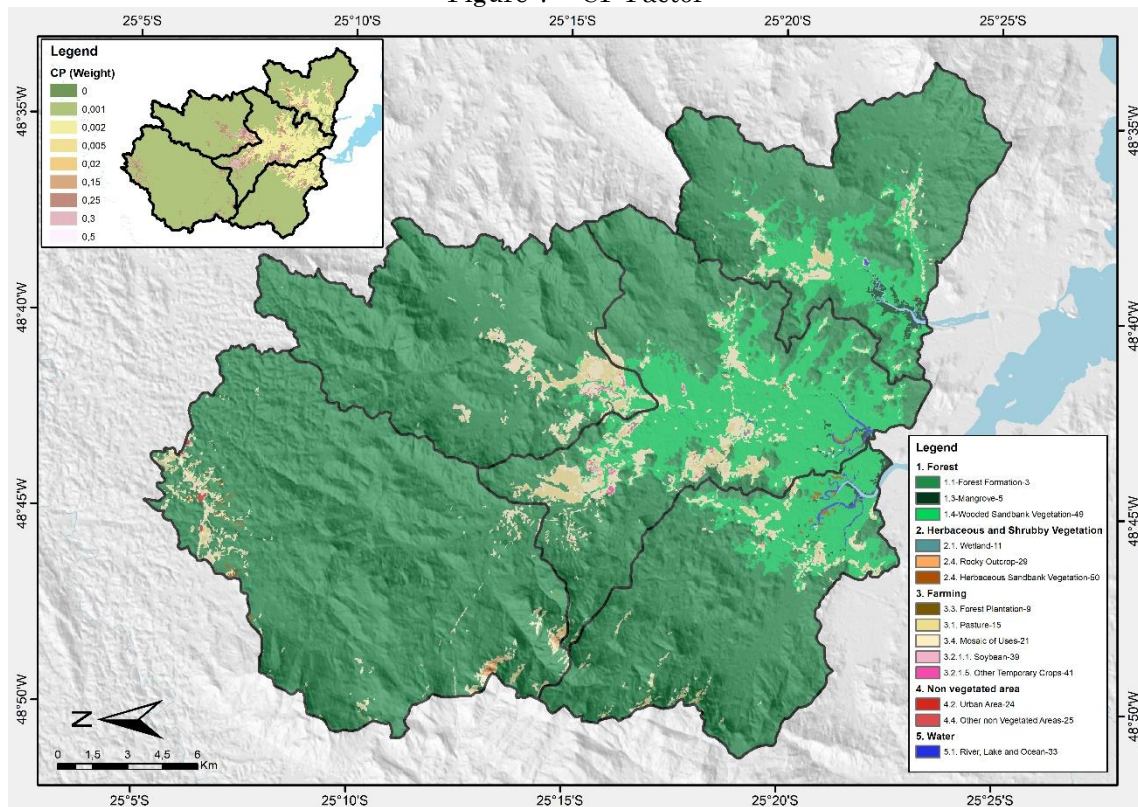
The CP factor represents the combined effect of land use and management (C) and conservation practices (P). The C factor expresses the relationship between soil loss under a given vegetation cover and management condition and the soil loss observed on bare ground, serving as an indicator of the protection provided by vegetation or surface residues. The P factor is associated with mechanical conservation practices—such as contour farming, terracing, and buffer strips—which reduce surface runoff velocity and enhance water infiltration.

The product of these two components results in the CP value, which ranges from 0 to 1. Values close to zero indicate high conservation efficiency and lower soil loss, whereas values near 1 reflect the absence of conservation measures. In this study, the CP factor (Table 3 and Figure 7) was assigned based on the land use and land cover mapped for the year 2024, with a 30-m spatial resolution ([Souza *et al.*, 2020](#)), following parameter values recommended in the literature for similar environmental conditions ([Miqueloni *et al.*, 2012](#); [Souza; Gasparetto, 2012](#)).

Table 3 – Land Use and Land Cover Classes and CP values

Code	Classes	C	P*	CP (C × P)	Final CP
1.1	Forest Formation	0,001 – 0,01	1,0	0,001 – 0,01	0,001
1.3	Mangrove	0,001	1,0	0,001	0,001
1.4	Flodable Forest	0,002	1,0	0,002	0,002
2.1	Wetland	0,005	1,0	0,005	0,005
2.4	Rocky Outcrop	0,000 – 0,001	1,0	~0,001	0
2.5	Herbaceous Sandbank Vegetation	0,02	1,0	0,02	0,02
3.1	Pasture	0,05 – 0,30	1,0	0,05 – 0,30	0,30
3.2.1.1	Soybean	0,20 – 0,30	1,0	0,20 – 0,30	0,30
3.2.1.5	Other Temporary Crops	0,25 – 0,35	1,0	0,25 – 0,35	0,30
3.3	Forest Plantation	0,10 – 0,20	1,0	0,10 – 0,20	0,15
3.4	Mosaic of Uses	0,15 – 0,30	1,0	0,15 – 0,30	0,25
4.2	Urban Area	0,00 – 0,05	1,0	0,00 – 0,05	0
4.4	Other non Vegetated Areas	0,30 – 1,00	1,0	0,30 – 1,00	0,5
5.1	River, Lake and Ocean	0,00	1,0	0,00	0

Source: Prepared by the authors (2025) *P was not calculated.

Figure 7 – CP FactorSource: Souza *et al.*, (2020).

Index of Connectivity (IC)

Since USLE calculates soil loss at the cell level within the raster matrices, it was necessary to incorporate complementary indices to characterize sediment connectivity dynamics in the study area. In this context, the Index of Connectivity (IC) (Borselli *et al.*, 2008; Cavalli *et al.*, 2013) was applied to estimate the efficiency

of sediment transfer from hillslopes to the drainage network (Equation 4).

$$IC = \log_{10} [D_{up}/D_{dn}]$$

Where Dup (potential contribution) is related to the contributing area and average slope of the hillside.

Ddn (resistance / distance to the channel) is the shortest path to drainage, weighted by slope and land use/cover.

Equation 5 was used to calculate the potential contribution (Dup):

$$D_{up} = \sum (A_i \tan(\theta_i))$$

Where A_i = accumulated contributing area in the cell.

θ_i = slope of the cell.

Equation 6 was used to calculate the channel distance variable - Ddn:

$$D_{dn} = \sum \left(\frac{d_i}{\tan(\theta_i) C_i} \right)$$

Where d_i = distance to the canal.

C_i = land cover factor (derived from the land use and occupation map).

Sediment Delivery Rate (SDR) and Effective Sediment Production Index (Effective)

The effective sediment export from the watershed was estimated by integrating the Universal Soil Loss Equation (USLE/RUSLE) with the Sediment Delivery Ratio (SDR). The SDR represents the fraction of eroded soil that effectively reaches the drainage network and is controlled by the degree of connectivity within the watershed (Borselli *et al.*, 2008; Vigiak *et al.*, 2012).

To estimate SDR in a spatially distributed manner, a logistic function was employed in which the Index of Connectivity (IC) acts as the explanatory variable, following approaches implemented in recent hydrosedimentological connectivity models (Cavalli *et al.*, 2013; López-Vicente *et al.*, 2013; Sharp *et al.*, 2020) (Equation 7):

$$SDR = \frac{1}{1 + e^{(a+b \times ic)}}$$

where a corresponds to the intercept of the logistic function

b to the slope coefficient that controls the slope of the curve. Reference values for a and b were proposed by Vigiak *et al.*, (2012) from calibrations in experimental basins in Europe, being $a \approx -0.56$ and $b \approx 0.17$. These parameters can be adjusted according to the local reality of each study.

The effective export of sediments (Effective) corresponds to the volume of material that is transferred from the slope to the drainage system. Its estimation considers the product of the potential soil loss, calculated using the Universal Soil Loss Equation (USLE), and the sediment delivery factor (SDR) (Equation 8):

$$E_{effective} = USLE \times SDR$$

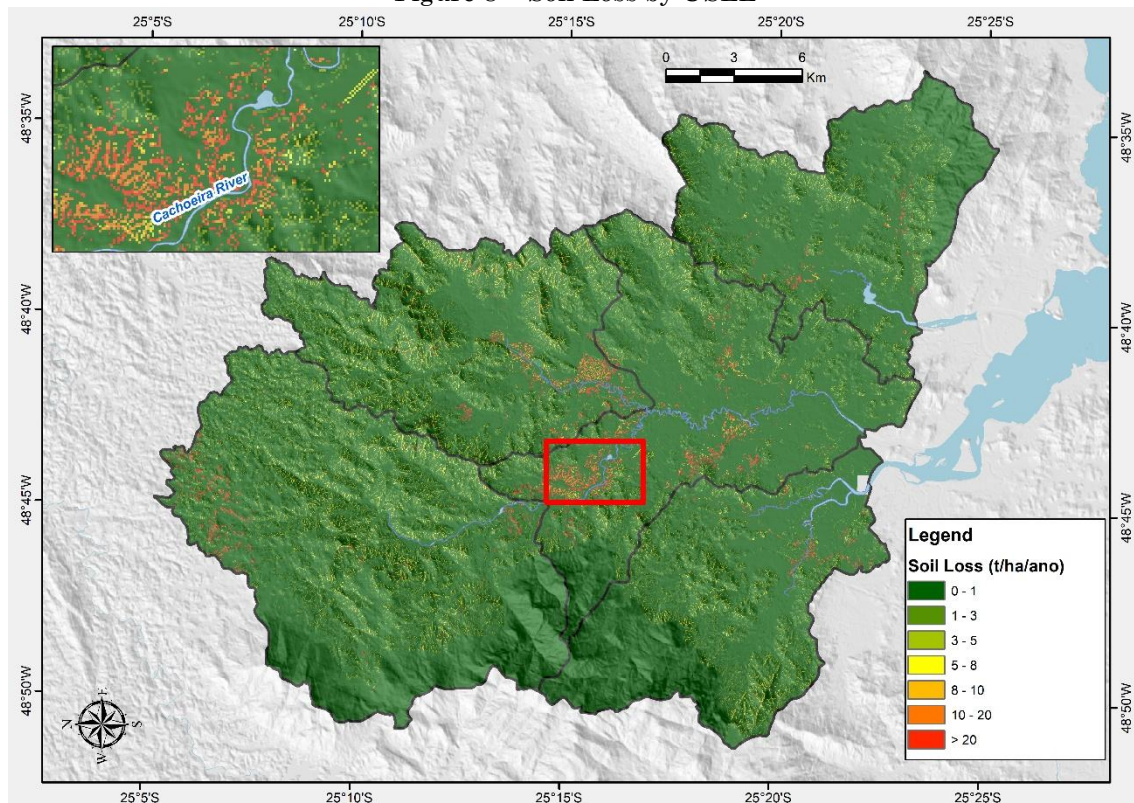
RESULTS

Soil loss by USLE

Areas most prone to soil loss are concentrated in environments with greater anthropogenic disturbance (Figure 8). Among the hydrographic basins, the highest concentration of elevated values occurs in the Lower Cachoeira basin. Significant sediment production also occurs in the Pequeno and Upper Cachoeira basins.

Average sediment production values are 1.44 t ha⁻¹ year⁻¹ in the Faisqueira basin, 2.32 t ha⁻¹ year⁻¹ in the Pequeno River basin, 2.85 t ha⁻¹ year⁻¹ in the Upper Cachoeira, 1.29 t ha⁻¹ year⁻¹ in the Cacatu basin, and 3.11 t ha⁻¹ year⁻¹ in the Lower Cachoeira basin.

Figure 8 – Soil Loss by USLE



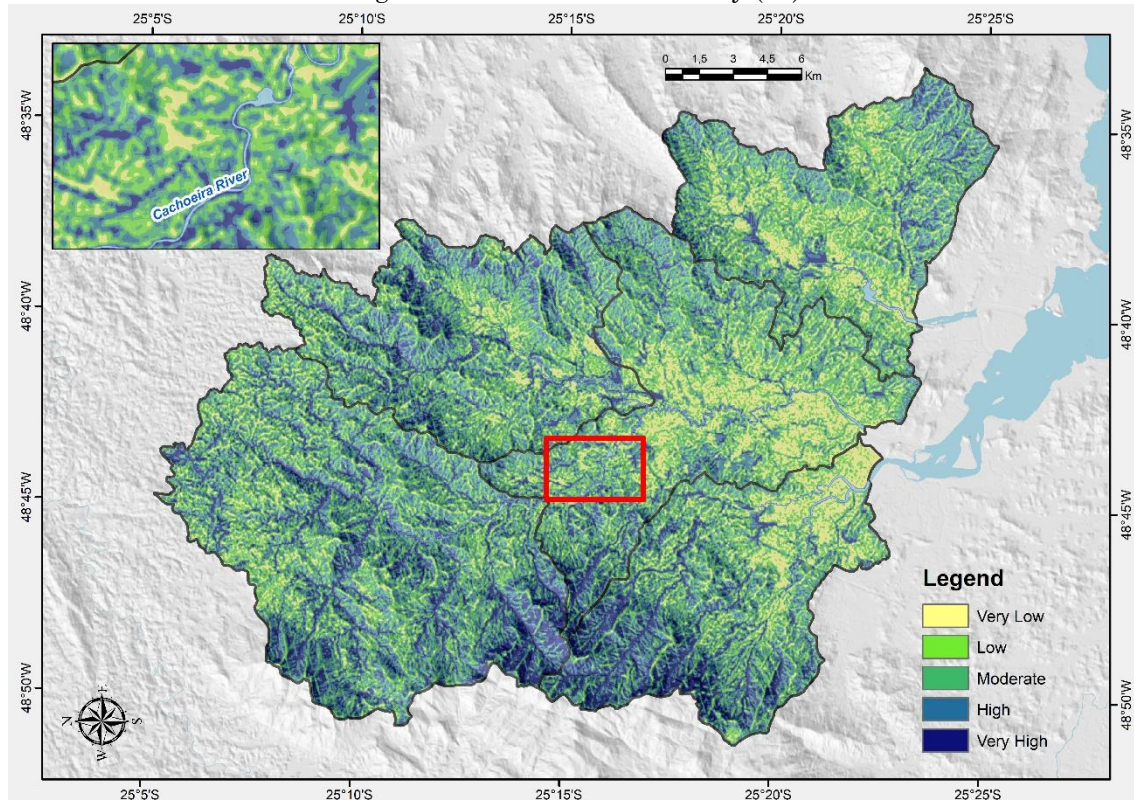
Source: The authors (2025).

Index of Connectivity

The Index of Connectivity (IC) values reflect the capacity of slopes to transfer sediments to the drainage network and are primarily controlled by slope gradient, proximity to channels, and vegetation cover density (Figure 9). Interfluvial zones and areas with gentler terrain generally exhibit very low to medium connectivity,

indicating a greater likelihood of sediment being retained before reaching the drainage network. In contrast, high and very high connectivity values are concentrated on the steepest slopes (greater than 15%) and along the main and secondary drainage channels, which act as preferential corridors for sediment transport toward watercourses.

Figure 9 – Index of Connectivity (IC)

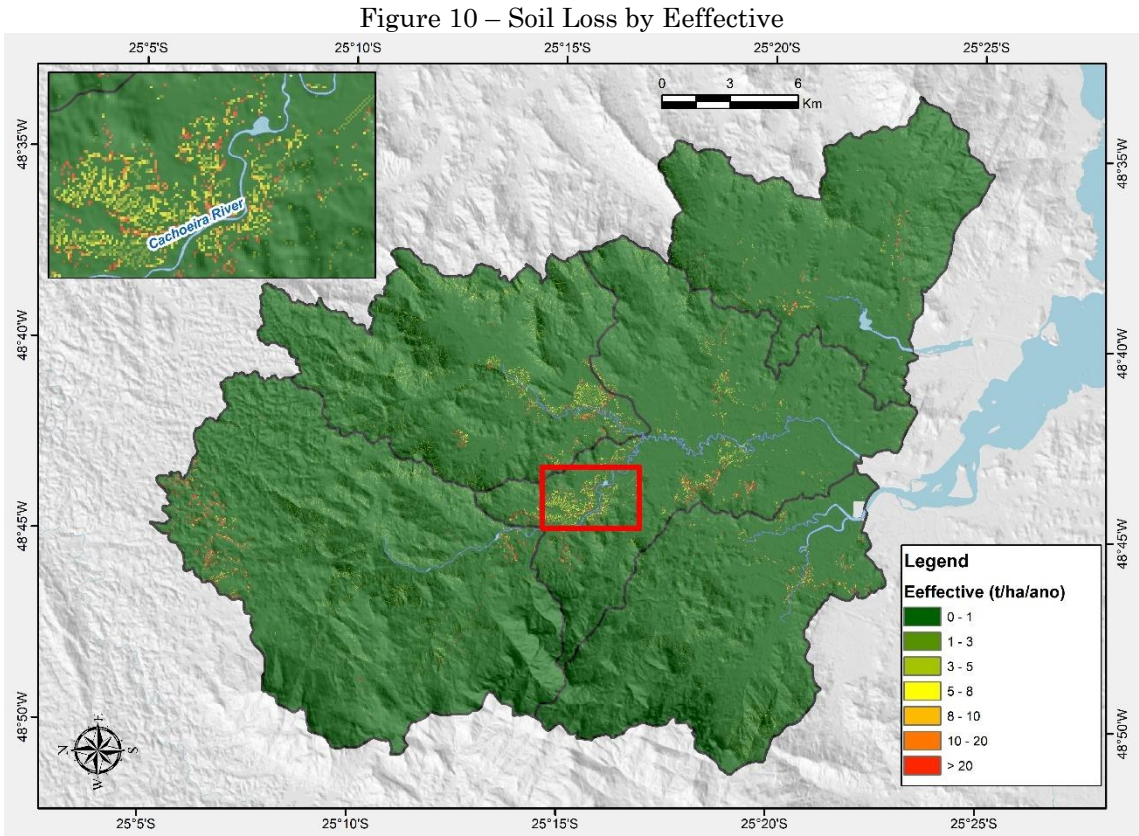


Source: The authors (2025).

Effective sediment export index (Effective)

The effective sediment export index ($\text{t ha}^{-1} \text{ year}^{-1}$) for the study area (Figure 10) provides an estimate of the spatial variability of soil loss that effectively reaches the drainage network. Most of the area exhibits low export values ($<1 \text{ t ha}^{-1} \text{ year}^{-1}$), indicating relative stability and a

reduced risk of sediment delivery to watercourses. However, certain sectors, particularly those adjacent to agricultural areas and along the Cachoeira River (highlighted in the figure 2e), show export rates exceeding $10 \text{ t ha}^{-1} \text{ year}^{-1}$. These zones represent critical areas, functioning as hotspots of soil degradation and high sediment-generation potential.



Source: The authors (2025).

DISCUSSION

The results indicate that the USLE tends to overestimate soil loss rates (Benavidez *et al.*, 2018; Alewell *et al.*, 2019; Meinen; Robinson, 2021), whereas its integration with the IC and SDR helps mitigate this limitation (Table 4), yielding spatial patterns that more closely reflect observed conditions (Souza *et al.*, 2025).

This variation is consistent with findings from tropical and temperate regions (Borselli *et al.*, 2008; Vigiak *et al.*, 2012). The occurrence of higher sediment production in highly connected areas reinforces the role of land use in shaping hydrosedimentological dynamics. Nevertheless, the combination of low altitudes, agricultural land use, and elevated soil loss in the Antonina Bay region underscores the environmental vulnerability of the system.

Table 4 – Soil losses estimated by USLE and Effective in the watersheds that drain into Antonina Bay

Catchment	Average soil loss USLE (t ha ⁻¹ ano ⁻¹)	Average soil loss Eefetivo (t ha ⁻¹ ano ⁻¹)	Reduction (%)
Faisqueira	1,44	0,41	71,52
Pequeno	2,32	0,66	71,55
Upper	2,85	0,83	70,87
Cachoeira	1,29	0,36	72,09
Lower	3,11	0,88	71,70
Cachoeira			
Total	2,02	0,62	71,54

Source: The authors (2025).

The results indicate average annual soil losses ranging from 1.29 to 3.11 t ha⁻¹ year⁻¹ according to the USLE, and from 0.36 to 0.88 t ha⁻¹ year⁻¹ for Eeffective, corresponding to

reductions greater than 70% when accounting for drainage connectivity. These values fall within the range reported for other watersheds in the Atlantic Forest (Fernandes *et al.*, 2014),

although they remain relatively high compared with areas under greater forest preservation.

In the Posses Stream watershed (state of Minas Gerais - southeast of Brazil), where extensive ecological restoration programs have been implemented, sediment production measured in experimental plots ranged from 0.01 to 0.05 t ha⁻¹ year⁻¹, indicating extremely low export rates (Gomes *et al.*, 2017). Conversely, in basins characterized by varying levels of Atlantic Forest cover (state of Rio de Janeiro - southeast of Brazil), sediment export monitored at the basin outlet ranged from 0.11 to 0.46 t ha⁻¹ year⁻¹, while USLE estimates reached up to 5.24 t ha⁻¹ year⁻¹, illustrating the method's tendency to overestimate potential erosion compared with measured sediment loads (Fernandes *et al.*, 2014).

At the continental scale, a European Union assessment indicated that only approximately 15% of predicted gross erosion reaches watercourses, with most sediments being trapped or deposited across the landscape (JRC, 2018). This pattern is widely documented: in basins with limited connectivity, effective export typically represents only a fraction of potential erosion. In the Ethiopian Rift, RUSLE-based estimates indicated erosion rates of 10–20 t ha⁻¹ year⁻¹, but only 21% of this material was delivered to the drainage network (Alemayehu *et al.*, 2025). In mountainous Mediterranean basins, observations highlight high hydrosedimentological variability and strong control of connectivity on sediment delivery (Francke *et al.*, 2018). Similarly, in southern Italy, rainfall–erosivity models indicated that only a small portion of generated material was exported, with most sediment retained within the basin (Diodato *et al.*, 2024).

The values obtained in this study, particularly those from the Eeffective model (0.4 to 0.8 t ha⁻¹ year⁻¹), are consistent with estimates reported for partially vegetated Atlantic Forest watersheds (Fernandes *et al.*, 2014), although they remain higher than those documented in areas under continuous forest cover (Gomes *et al.*, 2017). This contrast highlights the strong influence of land use and land cover on sediment export and reinforces the importance of integrating drainage connectivity into soil loss estimates.

The comparison between USLE and Eeffective demonstrates a marked reduction in soil loss values once connectivity is considered, producing results more consistent with those observed in other Atlantic Forest studies. Although based on different methodologies, the values obtained here also align with estimates generated by Rutyna *et al.*, (2021), who, using a

morphopedological and morphometric approach (Crepani *et al.*, 2001), reported average sediment production between 0.42 and 0.88 t ha⁻¹ year⁻¹ for the same basins draining into Antonina Bay. The similarity strengthens the consistency of the Eeffective index in capturing the effective export of sediments, mitigating the overestimation inherent to the standalone USLE (Benavidez *et al.*, 2018; Alewell *et al.*, 2019; Meinen; Robinson, 2021).

In the Antonina Bay region, land use emerges as the main factor driving sediment production, an effect also observed in other Brazilian biomes. Although not a central focus of this study, it is important to note that the Parigot de Souza Hydroelectric Plant, inaugurated in 1971, diverted part of the Capivari River flow into the Cachoeira River, increasing its annual discharge by approximately 33% (Soares; Santos, 2009). This change altered the hydrogeomorphological balance of the watershed and enhanced sediment export toward the coastal system, contributing to the silting of Antonina Bay.

Examples from other Brazilian biomes illustrate similar patterns. In the Cerrado, intensive agricultural expansion has increased sediment export (Oliveira; Leite *et al.*, 2018; Magalhães *et al.*, 2023), a trend also observed in the lower Cachoeira and Pequeno basins (Table 4; Figure 9). In the Amazon, deforestation for pasture increases hydrosedimentological connectivity (Barbosa; Fearnside, 2000), while in the Atlantic Forest, inadequate management on steep terrain intensifies erosion (Panachuki *et al.*, 2011). Despite regional particularities, interactions between relief, land use, and connectivity exhibit recurrent patterns, reinforcing the relevance of integrated methodologies such as the one applied in this study.

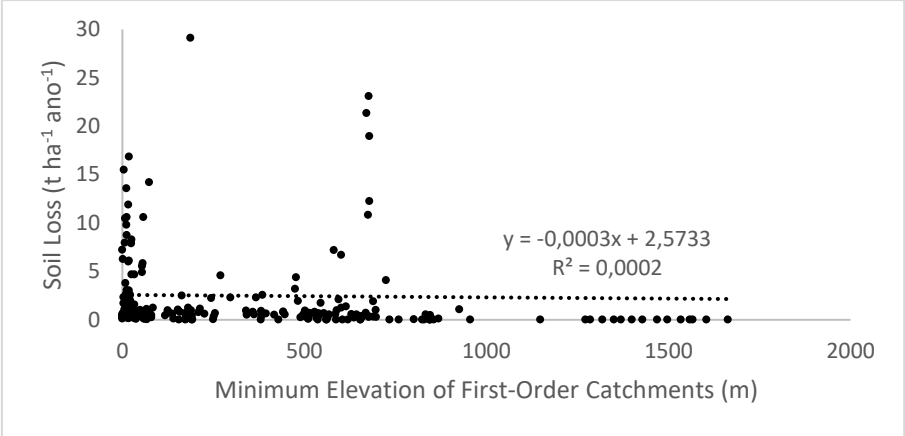
From a geomorphological perspective, areas with steeper slopes generally present greater susceptibility to sediment production. However, in the study area, these zones are predominantly covered by dense vegetation, which effectively inhibits erosive processes. Moreover, their greater distance from major floodplains limits the likelihood that eroded material will reach the main channels (Paula *et al.*, 2010). The Eeffective index corroborates this pattern, indicating that most sediments produced in these environments do not reach the drainage network.

To further illustrate these relationships, Figure 11 presents the association between the minimum elevation of first-order watersheds (m) and soil loss (t ha⁻¹ year⁻¹). The distribution of data points shows that minimum elevation

does not significantly explain the variation in estimated erosion, as indicated by the low R^2 value. Most watersheds with high losses ($>10 \text{ t ha}^{-1} \text{ year}^{-1}$) are located at elevations below 400 m. In contrast, above 600 m, soil loss values become sparse and remain close to zero. This pattern demonstrates that lower-lying areas near floodplains and Antonina Bay, where agricultural activities are concentrated, are more susceptible to elevated soil loss.

Conversely, basins situated at higher elevations, generally associated with better-preserved forest cover, and reduced anthropogenic pressure, exhibit minimal soil loss regardless of local topography. This contrast underscores the predominant role of land use and land cover in controlling erosion, outweighing the isolated influence of elevation.

Figure 11 – Graph illustrating the estimated soil loss by USLE x Elevation of the outlets



Source: The authors (2025).

When the estimated values are compared with field measurements (Souza *et al.*, 2025), a strong relationship emerges between land use and sediment production. Bare soil areas exhibit the highest losses ($0.035 \text{ t ha}^{-1} \text{ year}^{-1}$), a value approximately 25 times greater than that recorded under native forest cover ($0.0014 \text{ t ha}^{-1} \text{ year}^{-1}$). Conventional agriculture ($0.0064 \text{ t ha}^{-1} \text{ year}^{-1}$) and agroforestry systems ($0.0065 \text{ t ha}^{-1} \text{ year}^{-1}$) show intermediate values, markedly

lower than those observed for bare soil, yet still above the levels associated with native vegetation. These findings demonstrate that forest cover provides the most effective protection against erosion. They also reinforce that even the highest field-measured losses ($0.035 \text{ t ha}^{-1} \text{ year}^{-1}$ under exposed soil) remain lower than the estimates derived from USLE or from the Effective erosion values (Table 5).

Table 5 – Comparison of soil losses estimated by USLE and Effective in the watersheds that drain into Antonina Bay

Land Use	Sediment (g/plot/year) 2024*	Sediment (t ha ⁻¹ ano ⁻¹)
Agroforestry System	6,49	0,00649
Native Forest	1,41	0,00141
Agriculture	6,41	0,00641
Bare soil	35,03	0,03503

*Plots with 10 m²

Source: Souza *et al.*, (2025).

FINAL CONSIDERATIONS

The results show that the isolated application of the USLE substantially overestimates soil loss when compared with estimates obtained by integrating it with the Index of Connectivity (IC) and the Sediment Delivery Ratio (SDR).

Incorporating these parameters reduced potential erosion values by approximately 70%, indicating that most of the mobilized material is retained on slopes before reaching the drainage network. This finding directly highlights the importance of accounting for hydrosedimentological connectivity to more accurately represent actual sediment export.

The spatial analysis revealed that the highest soil-loss rates are concentrated in low-altitude areas with intensive agricultural use, particularly within alluvial plains, whereas higher-elevation regions covered by native forest exhibit minimal losses. The study therefore provides a detailed spatial assessment of the most vulnerable sectors, contributing to a clearer understanding of the geomorphological and anthropogenic controls on erosive dynamics. The resulting products, especially the IC and Eeffective maps, constitute valuable tools for environmental planning, enabling the identification of critical areas and guiding the allocation of priority management actions. Based on these results, the adoption of practices such as no-till agriculture, agroforestry systems, riparian forest maintenance, and improved land-use regulation in highly connected sectors is recommended, as these measures can significantly reduce sediment input to watercourses.

Finally, the methodology proved robust, compatible with readily accessible datasets, and replicable in humid tropical basins. The study directly advances erosion and sediment connectivity assessments, providing technical support for public policies related to water management, soil conservation, and the mitigation of siltation in environmentally sensitive regions.

ACKNOWLEDGMENT

We thank the Port Administration of Paranaguá and Antonina (APPA) for funding the project "Evaluation of soil loss and sediment production due to water erosion in experimental plots".

FOUNDING SOURCE

This article is derived from the project: Evaluation of soil loss and sediment production due to water erosion in experimental plots, funded by the Port Administration of Paranaguá and Antonina (APPA), SEI process number: 23075.062530/2021-14.

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José Guilherme de Oliveira: Conceptualization, Data curation, Data analysis, Investigation, Methodology, Software development, implementation and testing, Validation of data and experiments, Data presentation design, Writing – original draft, Writing – review & editing.

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ASSOCIATE EDITOR: Silvio Carlos Rodrigues. 

DATA AVAILABILITY: The data that support the findings of this study can be made available, upon reasonable request, from the corresponding author [José Guilherme de Oliveira].



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